

Crystal structure of a silica- and alkali-rich trioctahedral mica

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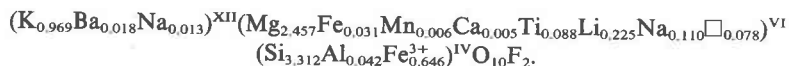
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Abstract

The crystal structure of a trioctahedral mica of unusual composition from melilite-bearing eruptive rocks has been determined ($R = 3.0\%$). This mica, which has excess silica ($\text{Si} > 3$) and alkalis ($\Sigma \text{Na} + \text{K} > 1$), is the first known layer silicate with octahedral sodium. In spite of the complexity of the octahedral layer, which includes Mg, Fe, Li, Ca, Mn, Na, Ti, and vacancies, there is no evidence for cation ordering between the two distinct octahedral sites. Both sites have the same mean M-O distance (2.077\AA) and the same electron density. The structural formula of the mica, based on chemical analysis and the structure refinement, is:



Introduction

Velde (1979) described a suite of trioctahedral micas of various compositions from melilite-bearing eruptive rocks. Of special interest is a group of silica-rich micas, characterized by high alkalis and ferric iron and low aluminum, in addition to significant amounts of Li, Ba, and Ti. These micas are unusual because $\Sigma(\text{K} + \text{Na} + \text{Ba}) \gg 1.0$, $\text{Si} \gg 3.0$, and $\Sigma(\text{Si} + \text{Al}) \ll 4.0$. Substantial deviations from the ideal biotite, $(\text{K},\text{Na})(\text{Mg},\text{Fe})_3\text{AlSi}_3\text{O}_{10}(\text{OH})_2$, therefore obtain. The principal objective of this study is to document the site distributions of cations in this mica.

Experimental

Euhedral single crystals of a silica-rich mica (sample Y253, Gragnani, 1972) from the type copaelite, at Cupaello, Italy, were selected for X-ray investigation. These crystals are unzoned and display a slight reverse pleochroism, characteristic of micas with tetrahedral Fe^{3+} (Hazen and Wones, 1972).

The composition of the crystals was determined by ion and electron microprobe analysis (Table 1). Fluorine was detected in significant amounts (>3 wt.%) by both ion and electron microprobe, but exact

quantitative determination could not be made. No OH^- was detected by infrared analysis and the mica thus has $F/(F + \text{OH}) \approx 1$. It is not known if there is a significant oxy-biotite component. Ion microprobe analysis indicated 0.75 ± 0.25 wt.% Li_2O (tourmaline standard), and boron in only trace amounts.

X-ray diffraction photographs of a crystal plate $140 \times 130 \times 50 \mu\text{m}$ revealed sharp diffraction maxima with none of the smearing or streaking of spots characteristic of deformed crystals or disordered stacking arrangements. Unit-cell parameters of this silica-rich mica (Table 2) are typical for a one-layer trioctahedral mica with potassium in the interlayer position (Hazen and Wones, 1972).

All X-ray reflections in one octant of reciprocal space ($\sin \theta/\lambda < 0.9$) were collected on an automated four-circle diffractometer with Nb-filtered $\text{MoK}\alpha$ radiation (Finger *et al.*, 1973). All reflections were corrected for specimen absorption ($\mu_1 = 21.0 \text{ cm}^{-1}$) as well as Lorentz and polarization effects. A total of 625 nonequivalent data were observed ($I > 2\sigma$). Trial refinements in space group $C2$, in which adjacent tetrahedra are symmetrically independent, and $C2/m$, in which all tetrahedra are equivalent, yielded similar results. The $C2/m$ refinement yielded the residual of

Table 1. Composition and calculated structural formula of silica-rich trioctahedral mica

	Wt. %	No. Atoms	Per Site		
			No. Atoms	No. Charges	No. Electrons*
SiO ₂	44.43	3.312	4.000	Tetrahedron +15.31	15.9e
Al ₂ O ₃	0.47	0.042			
B ₂ O ₃	trace				
Fe ₂ O ₃	11.38	0.646			
FeO†	0.49	0.031	2.922	Octahedra +5.68	11.5e
MgO	22.11	2.457			
MnO	0.09	0.006			
CaO	0.07	0.005			
TiO ₂	1.57	0.088			
Li ₂ O‡	0.75	0.225			
Na ₂ O	0.84	0.110	1.000	Alkali site +1.02	19.6e
		0.013			
BaO	0.62	0.018			
K ₂ O	10.18	0.969			
F§	4.5	2.000			
H ₂ O	(<0.1)				
Totals	97.50		7.922	+22.01	

*Number of electrons per site = Σ (number of atoms per site \times atomic number).

†Octahedral Fe²⁺/Fe³⁺ is unknown.

‡Based on ion probe analysis with tourmaline standard. Value cited is ± 0.25 weight percent.

§Fluorine present in significant amounts (>3%), based on ion and electron microprobe analysis. Calculations are based on end member fluoro-mica.

||Based on infrared spectra analysis by B. Velde.

Table 2. Unit-cell dimensions

Unit-Cell Parameter*	Value
a (Å)	5.3290(9)†
b (Å)	9.2300(15)
c (Å)	10.2191(11)
α (°)	90.00(1)
β (°)	99.98(1)
γ (°)	90.01(1)
V (Å ³)	495.0(1)

*Refinement of unit-cell parameters from diffractometer data was unconstrained (i.e. triclinic).

†Parenthesized figures represent esd's.

Table 3. Refinement conditions

Parameter	Value
Weighted R (%)*	3.0
R (%)†	3.0
No. of observations	625
Extinction parameter	0.6(+1.2) $\times 10^{-5}$

$$* \text{Weighted } R = [\Sigma w(|F_o| - |F_c|)^2 / \Sigma w F_o^2]^{1/2}$$

$$\dagger R = \Sigma ||F_o| - |F_c|| / \Sigma |F_o|$$

greatest significance, however, and the centrosymmetric space group, as observed in previous refinements of one-layer biotites, was thus selected. All observed reflections were included in the refinement of atomic positions, anisotropic thermal parameters, unconstrained occupancy factors for all metal sites and the fluorine site, and an isotropic extinction parameter. Conditions of refinement are recorded in Table 3, refined positional and thermal parameters in Table 4, and observed and calculated structure factors in Table 5. Least-squares weights were derived from

standard deviations based on counting statistics to which an amount equivalent to 2% of F^2 has been added.

Results and discussion

Interatomic distances and angles (Table 6) and the orientations and magnitudes of thermal vibration ellipsoids (Table 7) are similar to those of previously reported trioctahedral micas in the phlogopite-annite series (Hazen and Burnham, 1973; Bohlen *et al.*, 1980). The 5.7° tetrahedral rotation angle, which is

Table 4. Refined atomic coordinates, anisotropic temperature factors, equivalent isotropic temperature factors, and number of electrons per metal site in silica-rich trioctahedral mica

Atom	x	y	z	β_{11}	β_{22}	β_{33}	β_{12}	β_{13}	β_{23}	B Equiv.	No. of Electrons
T	0.5749(1)	0.1667(1)	0.2255(1)	0.0082(2)	0.0022(1)	0.0023(1)	-0.0000(1)	0.0007(1)	-0.0001(1)	0.87(1)	15.0(2)
M1	0	1/2	1/2	0.0063(5)	0.0012(2)	0.0026(2)	0	0.0008(2)	0	0.71(4)	11.7(2)
M2	0	0.8337(1)	1/2	0.0061(4)	0.0019(1)	0.0025(1)	0	0.0007(2)	0	0.78(3)	11.7(2)
K	0	0	0	0.0227(5)	0.0068(2)	0.0057(1)	0	0.0016(2)	0	2.40(3)	19.3(2)
O1	0.8198(3)	0.2355(2)	0.1672(2)	0.0236(8)	0.0097(3)	0.0042(2)	-0.0036(4)	0.0012(3)	-0.0003(2)	2.55(4)	
O2	0.5264(6)	0	0.1669(2)	0.0347(13)	0.0066(3)	0.0041(3)	0	0.0011(5)	0	2.61(6)	
O3	0.6298(3)	0.1670(2)	0.3901(1)	0.0096(5)	0.0026(2)	0.0027(1)	-0.0001(2)	0.0006(2)	0.0000(1)	1.03(3)	
F	0.1336(4)	0	0.3992(2)	0.0113(8)	0.0035(2)	0.0027(2)	0	0.0007(3)	0	1.19(5)	8.3(2)

*Number of electrons is refined for each cation site based on electron density. No constraints are used.

Table 6. Interatomic distances and bond angles

Bond	Distance (Å)	Atoms	Angle (°)	Bond	Distance (Å)	Atoms	Angle (°)
<u>Tetrahedron</u>				<u>M1 Octahedron</u>			
T-O1	1.653(2)*	01-T-01	108.16(7)	M1-O3[4]	2.093(1)	03-M1-O3[2]	94.80(8)
T-O1	1.654(2)	01-T-02	108.08(13)	M1-F[2]	2.044(2)	03-M1-O3[2]	85.20(8)
T-O2	1.655(1)	01-T-02	110.74(8)	Mean M1-O	2.077	03-M1-F[4]	96.04(5)
T-O3	1.657(2)	01-T-03	108.06(13)			03-M1-F[4]	83.96(5)
Mean T-O	1.655	01-T-03	110.75(9)	03-O3[2]	2.834(3)	Mean O-M-O	90.00
		02-T-03	110.95(10)	03-O3[2]	3.074(3)	(adjacent)	
			109.46	03-F[4]	2.767(2)		
O1-O1	2.678(1)			03-F[4]	3.075(2)	03-M1-O3	180.00
O1-O2	2.677(3)			Mean O-O	2.932	F-M1-F	180.00
O1-O2	2.677(3)			(adjacent)		(opposite)	
O1-O3	2.723(2)			<u>M2 Octahedron</u>			
O1-O3	2.725(2)			M2-O3[2]	2.092(2)	03-M2-O3	85.27(9)
O2-O3	2.729(3)			M2-O3[2]	2.094(1)	03-M2-O3[2]	85.26(6)
Mean O-O	2.701			M2-F[2]	2.044(2)	03-M2-O3[2]	94.49(6)
T-T[2]†	3.076(1)			Mean M2-O	2.077	03-M2-F[2]	83.93(7)
T-T	3.078(1)					03-M2-F[2]	96.33(7)
Mean T-T	3.077			03-O3[2]	2.835(3)	03-M2-F[2]	96.06(6)
<u>Interlayer Site</u>				03-O3	3.082(3)	F-M2-F	82.66(9)
K-O1[4]	3.022(2)			03-O3[2]	3.074(3)	Mean O-M2-O	90.00
K-O2[2]	3.020(3)			03-F[2]	2.767(2)	(adjacent)	
Mean K-O	3.021			03-F[2]	3.074(3)		
(inner)				03-F[2]	3.084(3)	03-M2-O3	179.66(9)
K-O1[4]	3.282(2)			F-F	2.707(4)	03-M2-F	177.83(7)
K-O2[2]	3.282(3)			Mean O-O	2.955	Mean O-M2-O	178.75
Mean K-O	3.282			(adjacent)		(opposite)	
(outer)							
Δ(K-O)	0.261						

*Parenthesized figures represent esd's.

†Bracketed figures represent multiplicity.

Table 7. Magnitudes and orientations of thermal ellipsoids

Atom	Axis	rms Displacement (Å)	Angle with respect to:		
			a (°)	b (°)	c (°)
T	r ₁	0.097(1)*	89(6)	6(5)	84(4)
	r ₂	0.107(1)	158(12)	87(6)	102(12)
	r ₃	0.110(1)	112(12)	95(4)	14(11)
M1	r ₁	0.071(5)	90	0	90
	r ₂	0.093(4)	178(7)	90	82(7)
	r ₃	0.115(4)	92(7)	90	8(7)
M2	r ₁	0.091(3)	90	0	90
	r ₂	0.092(3)	175(5)	90	85(5)
	r ₃	0.114(3)	95(5)	90	5(5)
K	r ₁	0.171(2)	73(9)	90	27(9)
	r ₂	0.172(2)	90	0	90
	r ₃	0.180(2)	17(9)	90	117(9)
O1	r ₁	0.145(3)	83(5)	83(4)	17(6)
	r ₂	0.168(3)	147(3)	121(3)	72(6)
	r ₃	0.218(3)	58(3)	148(3)	93(2)
O2	r ₁	0.146(5)	89(2)	90	11(2)
	r ₂	0.168(4)	90	0	90
	r ₃	0.223(4)	1(2)	90	101(2)
O3	r ₁	0.107(3)	83(17)	9(21)	86(14)
	r ₂	0.110(4)	139(17)	81(21)	120(17)
	r ₃	0.118(3)	50(16)	90(10)	150(17)
F	r ₁	0.118(4)	72(17)	90	27(17)
	r ₂	0.123(4)	90	0	90
	r ₃	0.128(4)	18(17)	90	117(17)

*Parenthesized figures represent esd's.

tron density of the T site is significantly greater than that of Si, Al or Mg. It is assumed, therefore, that all silicon plus aluminum, as well as 0.646 atoms of Fe³⁺, occupy this site. The total number of electrons on the tetrahedral site (15.1e⁻ based on refined electron density) and the mean T-O bond distance (1.655Å) from the structure refinement, conform well with this assumed tetrahedral site occupancy.

A second crystal-chemical constraint of all micas is that there can be no more than 1.0 interlayer cation per 11 oxygens. In this mica the sum of (K + Na + Ba) is 1.11, indicating that a significant fraction of these cations are not in the interlayer position. All K + Ba, as well as 0.013 atoms of Na, are assigned to a fully-occupied interlayer site. The total number of electrons (19.3e⁻) and the mean inner and outer K-O distances (3.021 and 3.282Å, respectively) from the structure refinement are consistent with this assignment.

All remaining cations must occupy the two octahedral sites. This necessary conclusion leads to one of the most complex site compositions in any known silicate. The resulting average octahedral layer compo-

sition, based on three octahedral cations per 11 oxygens, is:



The remarkable M1 and M2 sites thus have cations of valence 4+, 2+, and 1+, as well as vacancies, with cation radii ranging from 0.61 to 1.02 Å (Shannon, 1976). The total refined number of electrons on the two sites, M1 and M2, are both 11.7e⁻, and the mean M1-O and M2-O distances are both 2.077 Å. Although the complexity of the octahedral site chemistry precludes a unique solution to M1-M2 cation distribution, there is no evidence for cation ordering, consistent with previous refinements of both igneous (Hazen and Burnham, 1973) and metamorphic (Bohlen *et al.*, 1980) biotites, as well as synthetic BaLiMg₂AlSi₃O₁₀F₂ (McCauley and Newnham, 1973).

The resulting mica, though complex, may be represented as composed primarily (65%) of the ferriphlogopite end member, KMg₃Fe³⁺Si₃O₁₀F₂. Additional components include 23% of the magnesium lithium mica taeniolite, K(LiMg₂)Si₄O₁₀F₂, plus minor amounts of the phlogopite-annite series and a Ti-mica. Note that the calculated number of titanium atoms approximately equals the number of octahedral vacancies, as observed in many other biotites (Hazen and Burnham, 1973). The only previously unknown component of this mica is approximately 4% of an end member with Na in the octahedral layer.

Analytical uncertainties for fluorine, lithium, and octahedral Fe²⁺/Fe³⁺ result in small but significant uncertainties in the structural formula. For example, 10% oxy-biotite substituting for fluoro-biotite could increase the total number of cations from 7.92 to the ideal value of 8.00. Refined electron density of the fluorine site (8.3 e⁻) is smaller than the 9.0 e⁻ value expected for a pure fluoro-mica. This electron deficiency may imply the presence of an oxy-biotite component or undetected (OH)⁻. Uncertainties in lithium content may also affect the structural formula. If total lithium is 0.50 instead of 0.75 wt.%, then the num-

ber of octahedral cations will decrease from 2.92 to 2.85.

This unusual mica serves a special role in the crystallization history of the melilite-bearing eruptive rocks in which it is found. It appears to concentrate alkali metals, including potassium, sodium, and lithium, in an environment that is deficient in aluminum and thus underscores the adaptability of the mica structure to a wide range of compositional limits.

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